Generation of solitary waves in a pycnocline by an internal wave beam

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Acronyms: IW(B) = internal wave (beam), ISW = internal solitary wave, Φ -speed = phase speed.

Introduction

Internal solitary waves have been detected in the Bay of Biscay far from the coast, at a distance too large for the waves to have been generated at the **shelf break** (cf. [1], [2], [3], [4]). Proposed mechanism for their presence far from the coast ([1]): internal wave **beam** (here the internal tide) **hitting the seasonal thermocline** ⇒ large interfacial displacement \Rightarrow trains of ISWs. Such a mechanism is rare (western) Europe, southwestern Indian ocean) ⇒ very restrictive conditions should hold for it to occur.

Mode-1 solitary waves



- > Exp. E1: $\Delta_p = 2\% \Rightarrow$ interfacial Φ -speed in a 2-layer fluid with same density jump: $c^*=6.3$ cm.s⁻¹ $> \lambda_{x} \approx \Lambda = 60$ cm, horiz. Φ -speed of the IWB: $v_{d} = 4.1$ cm.s⁻¹. > In [5], $\gamma = c^*/N_{\rho}H$ should be 0.12; here: $\gamma = 0.11$. > In [6], $\alpha = N_0 \lambda / \pi c^*$ should be 1; here: $\alpha = 1.3$. In one forcing period, 3 stages can be distinguished: **1.** impact of the IWB \Rightarrow partial reflection, **partial** transmission in the pycnocline
- 2. **nonlinear steepening** of the transmitted part



Theoretical works ([5], [6]): the value of the density jump across the pycnocline has to be of moderate **strength** so that the Φ -speeds of the interfacial wave and of the forcing IWB have the same order of magnitude.



* ISW packets in the Bay of Biscay (SAR images), traveling away from the shelf break. **Packets close to the shelf break (blue)** are to be distinguished from those emerging far from it (red). From [4].

This mechanism has never been observed nor simulated numerically up to now. We present non-rotating, direct numerical simulations which **confirm this mechanism**. We show that various modes of ISW can be generated, whereof occurrence can be controlled.

Numerical Set-up

Reproduction of the geometry and fluid parameters of the Coriolis turntable

* IW field (snapshot). Top: isopycnals around

the pycnocline, a train of 3 ISWs is indicated. Bottom: horizontal velocities, whole field.

 \Rightarrow IW trapped in the upper layer,

3. nonhydrostatic disintegration into 3 ISWs. Partial transmission back in the lower layer \Rightarrow scattering of the IWB (*cf.* [5], [7]).

* Top: space-time representation of the vertical displacement of the pycnocline. Bottom: temporal Fourier analysis of top.

Mode-n solitary waves: a modal resonance condition

 \succ Taylor-Goldstein equation for frequency $\Omega \Rightarrow$ IW modes of frequency Ω $\frac{d^2 W_n}{2} + \frac{N^2(z) - \Omega^2}{2} W_n = 0$ $C_n^2(\Omega)$ W: vertical structure of nth mode c_n(Ω): Φ-speed of n^{th} mode > When $\Omega > N_{\rho}$ (trapped waves): only small variations of c_{ρ} with $\Omega \Rightarrow c_{\rho}(\Omega) \approx \hat{c}_{\rho}(\Omega)$ **Resonance condition to generate** mode-*n* ISWs: $v_{\perp} \sim \hat{c}_{\perp}$ \rightarrow Plot of the Φ -speeds $\Box_1(\Omega) / v_{\phi}$ (scaled by v_{A}) versus ΩC₂(Ω) / V_φ (scaled by the forcing _-3

Illustration of the resonance condition for mode-2 ISWs. (Experiment E2) $\succ \Delta_{\mu} = 3.4\%, \lambda_{\chi} \approx 26 \text{ cm}, v_{\mu} = 1.8 \text{ cm}.\text{s}^{-1}.$

 $\Box_1^{(\Omega)} / v_{\phi}$

......c₂(Ω) / v_φ

____c₃(Ω) / ν_φ

→ Plot of the Φ-speeds: now $V_{\phi} \sim \hat{C}_{2}$ \$ Developed wave field & zoom on the pycnocline (top)



Illustration of the resonance condition for mode-3 ISWs. (Experiment E3) $\succ \Delta_{p} = 4\%, \lambda_{\gamma} \approx 15 \text{ cm}, v_{\phi} = 1 \text{ cm}.\text{s}^{-1}.$





- (Grenoble), where experiments on this subject have been performed in 2008.
- > **MITgcm** code: incompressible nonhydrostatic Boussinesq equations, centered 2nd-order finite volume scheme.
- Continuous temperature profile (see figure, red), idealized version of the summer stratification in the Bay of Biscay.
- Forcing: temporally oscillating velocity field at the left boundary (see figure,). Sponge layer on the right boundary. magent
- Explicit scheme, no parametrization, nothing "under the hood".
- > Parameters that will be varied: Δ_p (see fig., red), Λ (see fig., magenta), L and

resolution in **3 experiments**: E1, E2 and E3, in which mode-1, -2 and -3 ISWs develop respectively.





Mode-n (n≥2) **solitary waves: a Bragg-like resonance condition**



- * Magnification of the velocity fields at the beam impact in the pycnocline for E2 (left) and E3 (right)
- Careful observation of the IWB transmission in the pycnocline reveals refraction and modification of its vertical structure.
- > Simple model, approximations:
- 1. Stratification → 3 layers (see →)
- 2. IWB → plane wave
- 3. WKB-like Ansatz (rays)







> For E2: $\mu_2 = 0.88$; for E3: $\mu_3 = 0.95$

⇒ excellent agreement of this model with the numerics





parameters used in the text.

~ 1 m

2 cm

1 cm

45°

70

0.4, 4 mm

10⁻⁷ m.s⁻²

Conclusions and perspectives

References

[1] New, A. L. & Pingree, R. D. 1990. *Deep-Sea Res.*, **37**, 513–524. [2] Pingree, R. D. & New, A. L., 1991. J. Phys. Oceanogr., **21**, 28-39 [3] New, A. L. & Pingree, R. D. 1992. *Deep-Sea Res.*, **39**, 1521-1534. [4] New, A. L. & Da Silva, J. C. B. 2002. *Deep-Sea Res. (I)*, **49**, 915–934. [5] Gerkema, T. 2001. J. Mar. Res. 59, 227–255. [6] Akylas, T. R., Grimshaw, R. H. J., Clarke, S. R. & Tabaei, A. 2007. J. Fluid Mech. 593, 297–313. [7] Mathur, M. and Peacock, T. 2009, J. Fluid Mech., 639, 133-152. [8] Grisouard, N., Staquet, S. & Gerkema, T., 2010. Submitted to J. Fluid Mech.

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Summary

First numerical evidence of the generation of internal solitary waves by an internal wave beam.

\succ Possibility of generating ISWs of any mode.

> Two different resonance conditions to select the

mode, both show that the slower the phase speed of the beam, the higher the ISW mode.

Free surface

simplified density profile

original density profile 1

Perspectives

Validation against experiments (ongoing). Application to realistic simulations (ongoing). Integration of realistic effects (shear flow, background IW) field...).

Validation against in situ measurements.